

## Tectonic modulation of caldera volcanism on the South Aegean Volcanic Arc

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## ABSTRACT

Many highly hazardous, caldera-forming explosive eruptions occur in extensional tectonic regimes, but the role of lithospheric rifting in modulating caldera volcanism remains enigmatic. IODP Expedition 398 deep-drilled the volcano-sedimentary infills of submarine half-grabens around Santorini caldera on the continental South Aegean Volcanic Arc. Here we use the volcanic tephra archives to produce a high-resolution eruptive chronostratigraphy for Santorini, to ground-truth seismic stratigraphy, and to extract an integrated timeline of volcano-tectonic couplings. The rift basins contain several submarine volcanoclastic megabeds from the caldera-forming eruptions of Santorini and one from the Kos caldera. The thickest megabed succession is < 250,000 yrs old and lies on a seismic reflection onlap surface that records a phase of rapid rifting. Sedimentation lagged behind subsidence during this rifting phase, creating bathymetric troughs. Integrating submarine core-seismic and onland datasets, we propose that rifting may have driven the transition of Santorini from a prolonged state of effusive and minor explosive activity (~550 – 250 ka) typical of arc stratovolcanoes to one of repeated caldera-forming eruptions (<250 ka). Rapid rifting may have amplified the normal internal dynamics of the magmatic system in three ways, driving the volcano into a sustained, highly explosive state: (1) an increase in the supply of mantle-derived basalt, (2) enhanced shearing, permeability, and melt percolation in the transcrustal magmatic system, and (3) the development of horizontally extensive magma reservoirs. Broadly simultaneous transitions into caldera-forming activity of the widely separated Santorini and Kos Volcanoes suggest that the two magmatic systems are linked by plate-scale lithospheric stresses.

## 1. Introduction

Caldera-forming explosive eruptions of volcanoes discharge large volumes of silicic magma and impact humans, infrastructures and environments on local to global scales (Newhall et al., 2018; Cassidy and Mani, 2022). Several decades of research have led to models of the internal controls of caldera volcanism. The mantle-derived basalt that drives caldera systems generates silicic magmas by fractional crystallisation, crustal melting, and pluton defrosting in transcrustal mush systems (Bachmann and Huber, 2016; Cashman et al., 2017; Sparks et al., 2019). Caldera-forming eruptions occur when the fluxes of melts, fluids and heat into the crust are sufficiently high to raise the brittle-ductile transition and create the crustal rheological conditions necessary for the growth and maturation of large magma reservoirs (Jellinek and DePaolo, 2003; Degruyter et al., 2016; Townsend et al., 2019). These reservoirs can then trap the ascending melts in a thermo-mechanical runaway, leading to caldera-forming eruptions (Karakas and Dufek, 2015; Bouvet de Maisonneuve et al., 2021).

However, the internal processes of sub-caldera magmatic systems can be influenced by lithospheric tectonics (here used to describe the upper plate strain in a subduction-driven system; Fig. S1). This is particularly true in extensional environments such as continental rift zones (e.g., East African Rift, Hutchison et al., 2016; Taupō Volcanic Zone, Wilson et al., 2009; South Aegean Volcanic Arc, Hooft et al., 2017). Many calderas lie in extensional settings, suggesting that rifting helps promote mantle inputs high enough to push the magmatic systems over the threshold into the caldera-forming state (Hughes and Mahood, 2008; Burns et al., 2015). However, tectonic strain and internal magmatic processes at caldera systems operate on variable timescales, from  $10^0$  yr for eruption triggering to  $10^6$ – $10^7$  yr for large ignimbrite flare-ups, making it challenging to establish causal relationships (Gravley et al., 2016; Sternai, 2020; Berryman et al., 2022; Muirhead et al., 2022; Brune et al., 2023). While the effects of lithospheric extension on the locations of rift-hosted volcanoes, and the feedbacks between magmatism and crustal stresses, have been widely investigated (Rowland et al., 2010; Brune et al., 2023), the effects of rifting on volcanic eruption style and magnitude at caldera systems are, with some exceptions (Wilson et al., 2009; Gravley et al., 2016; Hutchison et al., 2016), less well understood.

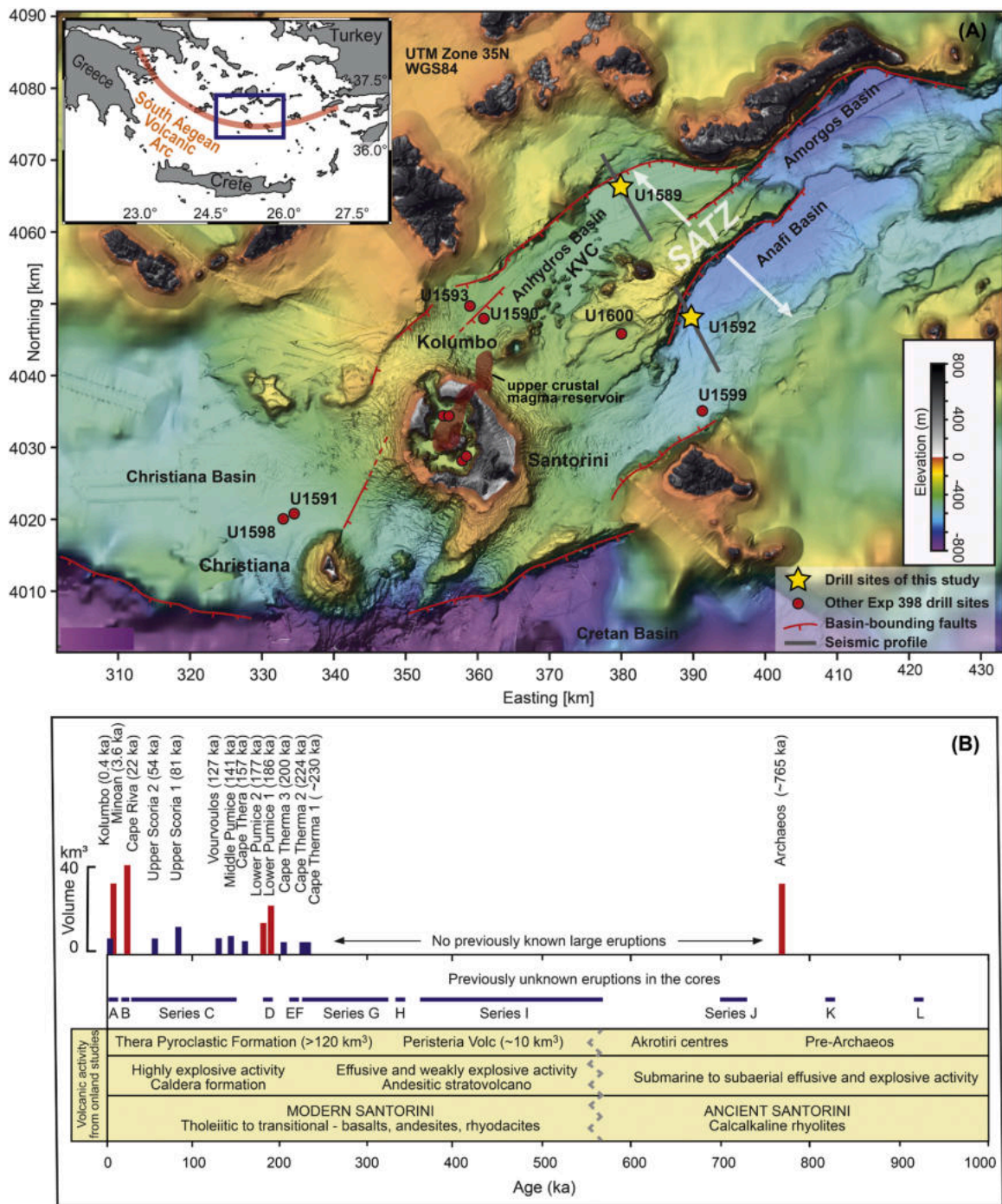
Here we investigate the inter-relationships between lithospheric rifting and caldera volcanism at Santorini Volcano (South Aegean Volcanic Arc; Fig. 1a) on the  $10^4$  –  $10^5$  yr timescales characteristic of arc volcano caldera cycles. The volcanic history of Santorini has been partly unraveled by studies of marine tephrostratigraphy and on-island volcanic products (Fig. 1b; Druitt et al., 1999; Wulf et al., 2020; Kutterolf

et al., 2021), and the offshore seismic stratigraphy of the surrounding rift basins has been documented and interpreted (Hübscher et al., 2015; Nomikou et al., 2016, 2018; Tsampouraki-Kraounaki et al., 2021; Preine et al., 2022a,b). A close relationship between faults, dyke swarms, and volcanic vents with structural lineaments at and around Santorini has been shown to reflect a crustal tectonic influence (Kokkalas and Aydin, 2012; Feuillet, 2013; Hooft et al., 2017; Tsampouraki-Kraounaki et al., 2021; Preine et al., 2022a,b; Drymoni et al., 2022), and analysis of the distribution and orientation of faults and fractures from seismic tomography data suggests a fundamental influence of regional-scale stresses (Heath et al., 2021). Preine et al. (2022a) proposed from interpretation of seismic stratigraphy that a period of caldera-forming explosive volcanism on Santorini was preceded by a phase of rapid rift extension. Testing this hypothesis requires well-dated, high-resolution records of both volcanic and tectonic events that can be precisely correlated. IODP (International Ocean Discovery Program) Expedition 398 deep-drilled the volcano-sedimentary infills of marine rift basins around Santorini (Druitt et al., 2024b). Explosive eruptions shower ash and/or pour pyroclastic currents into the marine basins, generating a rich archive of volcanic events that is potentially more complete, and extends further back in time, than that recorded onland (Fig. 2). Using biostratigraphic and tephrostratigraphic age constraints, and integrating the published seismic stratigraphy, we investigate a volcano-tectonic link at Santorini and discuss possible mechanisms of couplings and feedbacks between lithospheric extension and caldera volcanism on the South Aegean Volcanic Arc.

## 2. The volcano-tectonic setting

The volcanoes of Santorini and the adjacent submarine Kolumbo volcanic chain (Kolumbo seamount and >20 smaller eruptive centres) lie on ~25 km of stretched continental crust (Hooft et al., 2017) (Fig. 1a). Along with the extinct Christiana Volcano, they form the Christiana-Santorini-Kolumbo Volcanic Field, one of five volcanic fields on the South Aegean Volcanic Arc (Vougioukalakis et al., 2019). Christiana was active in the Pliocene through to the Early Pleistocene, but has been extinct since ~1.6 Ma (Preine et al., 2022b). Santorini has had two main phases of development (Fig. 1b):

- **Ancient Santorini (>550 ka)**, also known as ‘old’ or ‘ancestral’ Santorini, erupted calcalkaline rhyolites from shallow-submarine to subaerial vents between >650 and 550 ka (the onland ‘Early Centres of Akrotiri’ of Druitt et al., 1999). Druitt et al. (2024a) also identified a voluminous (Dense Rock Equivalent volume >30 km<sup>3</sup>) submarine rhyolitic pumice deposit, the Archaeos Tuff, from a mid-Pleistocene



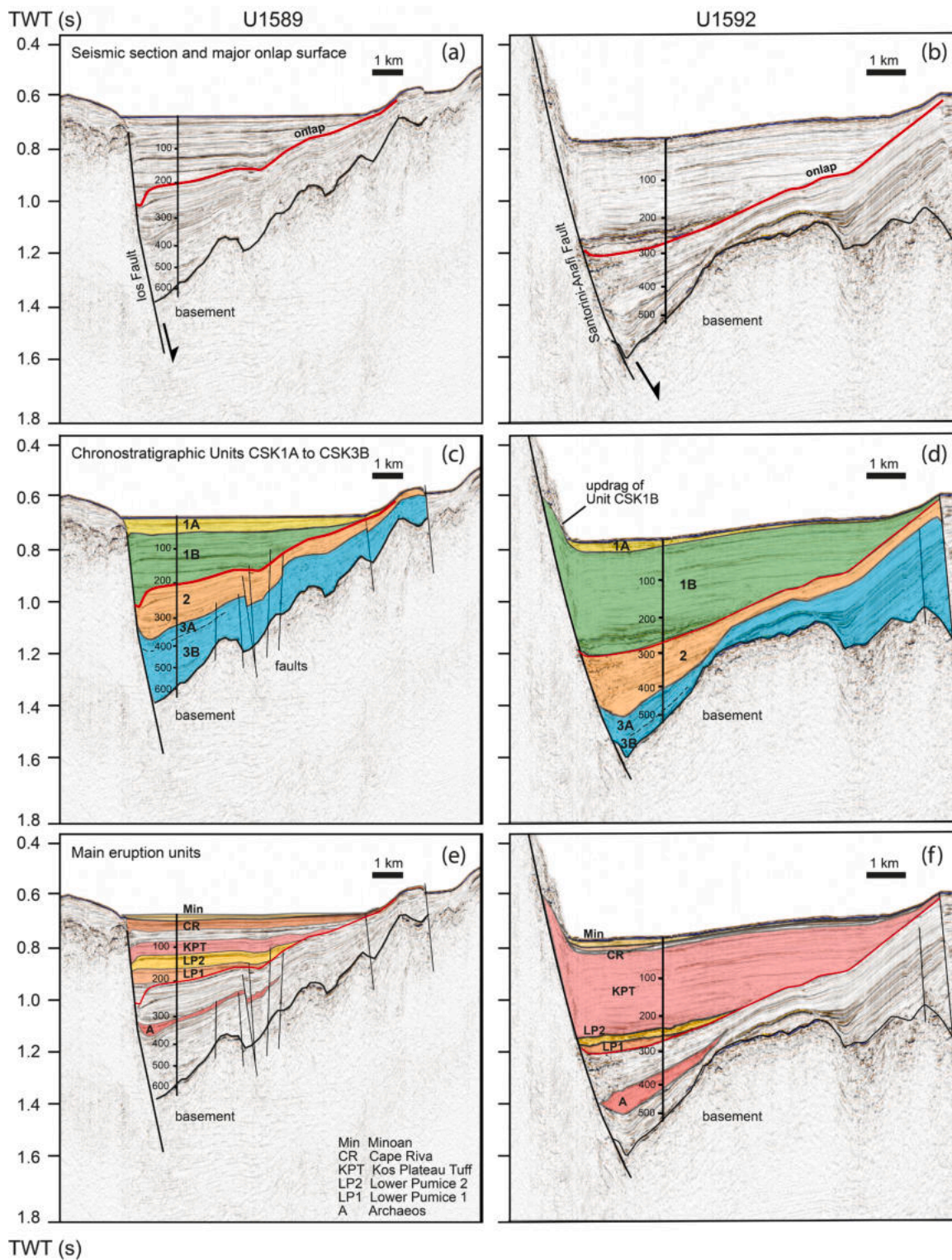
**Fig. 1.** (a) Location of the Christiana-Santorini-Kolumbo Volcanic Field and its relationship to the Santorini-Amorgos Tectonic Zone (SATZ) on the South Aegean Volcanic Arc (inset: blue rectangle) and of the IODP drilling sites outside of Santorini caldera (red dots, with site numbers). The two sites most relevant to the present paper (U1589 and U1592) are shown as yellow stars. Basin-bounding faults are shown in red, with ticks on the down-thrown side. The semi-transparent brown areas marks the upper crustal storage zone imaged by McVey et al. (2019). KVC: Kolumbo Volcanic Chain. See Methods for sources of bathymetric data. Modified from Druitt et al., (2024a). (b) Summary of the explosive volcanic history of Santorini Volcano, showing the main eruptions that produced the volcanic layers found in the IODP cores. The named eruptions known previously from onland studies (Druitt et al., 1999) are shown as vertical bars, the heights of which are proportional to eruptive volume (dense-rock equivalent (Kutterolf et al., 2021; Druitt et al., 2024a)). Large caldera-forming silicic eruptions (CFSE in the main text) are shown as red bars, and smaller explosive eruptions as blue bars. Eruptions relating to previously unknown volcanic layers discovered in the IODP cores of the present paper are shown as blue horizontal lines and labelled by a letter corresponding to the name given either to a given tephra layer or to a series of tephra layers. The yellow boxes show the corresponding timeline of volcanic activity known previously from onland studies.

explosive eruption of Ancient Santorini in cores from IODP Expedition 398. The age of the Archaeos eruption was estimated as 520 ka using biostratigraphic data, but <sup>40</sup>Ar/<sup>39</sup>Ar plagioclase and U-Pb zircon dating now place it at ~765 ka (S. Beethe and A.K. Schmitt, unpublished data).

- **Modern Santorini (<550 ka)** has erupted basalts to rhyodacites that are tholeiitic to transitional tholeiitic-calkalkaline in character. Its history can be divided into two broad stages. For the first ~300 kyr, Modern Santorini exhibited the effusive and weakly explosive behavior typical of andesitic stratovolcanoes (including the <10

km<sup>3</sup>, 530 – 430 ka ‘Peristeria’ stratocone; [Druitt et al., 1999](#)). It then transitioned at ~250 ka into a series of highly explosive explosions that discharged >120 km<sup>3</sup> of intermediate to silicic magma and laid down the ‘Thera Pyroclastic Formation’ (TPF) ([Druitt et al., 1999](#); [Kutterolf et al., 2021](#)). At least four of the TPF eruptions formed calderas, grouped into two time packages: 186 – 177 ka (Lower

Pumice 1 and Lower Pumice 2) and 22 – 3.6 ka (Cape Riva and Minoan). The TPF stage of Modern Santorini is characterized by (1) a concentration of highly explosive activity, (2) repeated Caldera-Forming Silicic Eruptions (CFSEs), and (3) a high volcanic output ([Fig. 1b](#)).



**Fig. 2.** (a-b) Seismic profiles for the Anhydros and Anafi basins with the major onlap surface shown in red. (c-d) Chronostratigraphic units, described in the text. (e-f) The caldera-forming silicic eruptions (CFSE in the main text) of Santorini and the Kos Plateau Tuff from the Kos Volcano. Core-seismic integration from [Preine et al. \(2025\)](#).

Kolumbo Volcano last erupted explosively in 1650 CE, and seismic profiles across it reveal multiple, stacked edifices from older eruptions (Hübscher et al., 2015).

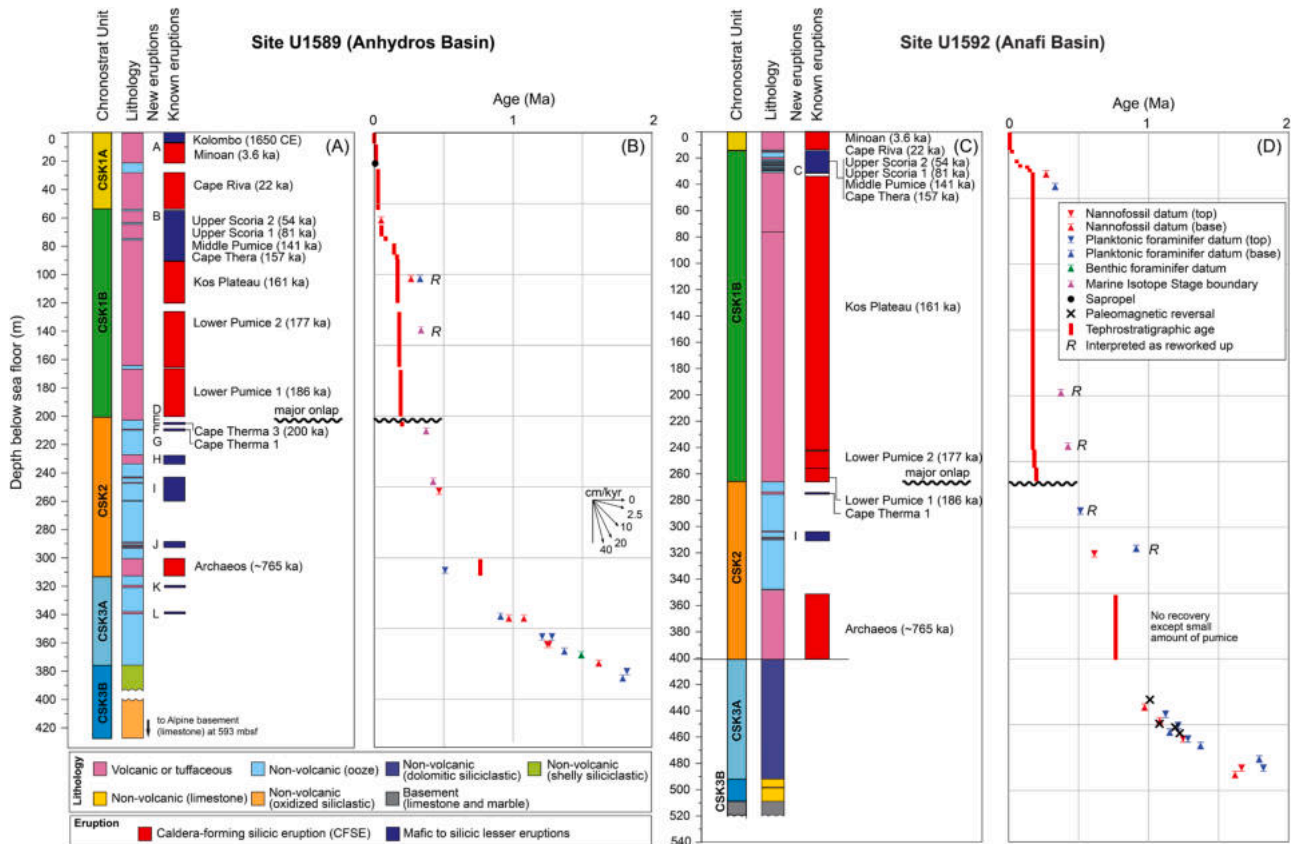
Santorini lies at the southern end of a NE-SW system of en-echelon continental rift basins (Santorini-Amorgos Tectonic Zone; SATZ) that cut the volcanic arc and extend into the back-arc region (Nomikou et al., 2016, 2018; Hooft et al., 2017; Tsampouraki-Kraounaki et al., 2021). Extension of the SATZ is driven by slab rollback and lateral bending, facilitated by slab tears underneath the Anatolian plate and eastern Aegean (Jolivet et al., 2013; Sternai et al., 2014). The possible existence of a right-lateral oblique component to the faulting is debated (Tsampouraki-Kraounaki et al., 2021; Crutchley et al., 2023). Three half-graben basins (Anhydros, Anafi, Amorgos) defined by multiple overstepping normal faults and relay ramps constitute the SATZ (Hooft et al., 2017; Crutchley et al., 2023) (Figs. 1a and 2). While the western Anhydros Basin contains the Kolumbo Chain volcanoes, the Anafi Basin and eastern Anhydros Basin are non-volcanic (Hübscher et al., 2015; Nomikou et al., 2016; Preine et al., 2022a). Seismic tomography has imaged a NE-SW elongated upper-crustal reservoir of melt-bearing magmatic crystal mush that extends from Santorini towards Kolumbo (Fig. 1a; McVey et al., 2019). Other melt storage regions have been imaged seismically NE of Santorini at mid-crustal depths (Autumn et al., 2025; Hufstetler et al., 2025).

A dense array of single and multichannel seismic profiles across the SATZ reveals rift sediment fills up to 1.4 km thick, with multiple seismic stratigraphic packages (Hübscher et al., 2015; Nomikou et al., 2016, 2018; Preine et al., 2022a, b). The SATZ is cut to the south by the earlier, ENE-WSW Christiana Basin of Pliocene age (Heath et al., 2019; Preine

et al., 2022b) (Fig. 1a).

### 3. Materials and methods

IODP Expedition 398 deep-drilled eight sites outside of Santorini caldera on the R/V JOIDES Resolution (Fig. 1a). We here focus on Site U1589 in the Anhydros Basin and Site U1592 in the Anafi Basin, each of which hosts a different sediment distributary branch of the volcanic field. The drill cores were logged on board (Fig. 3; Table S1), taking into account artefacts of drilling and core recovery such as sediment mixing, shear-induced uparching, brecciation, biscuiting and ash liquefaction (Jutzler et al. 2025). Physical properties including sediment density and P-wave velocity were measured *in situ* on the cores. Samples of the volcanic layers were collected for chemical analysis of glass shards for chemical fingerprinting and correlation purposes: major elements by electron microprobe and trace elements by laser-ablation ICPMS (Table S2). Correlations used previously published databases of glass chemistry from studies of Santorini and Kolumbo tephra in shallow gravity cores (Kutterolf et al., 2021). Descriptions of volcanic layers from previously known eruptions (Table S3) are given in Table S4, and those of the newly recognized eruptions are given in Table S5. Ages of some volcanic layers were determined by chemical correlation with marine tephra layers from published studies of shallow marine gravity cores, the ages of which are already known (Table S3; Wulf et al., 2020; Kutterolf et al., 2021). Biostratigraphic ages of the sediments were determined ship-board using assemblages of foraminifera and calcareous nannofossils, based on one sample per 9.5 m or 4.7 m core (Tables S6 and S7). Sapropels of previously known age were recognized using the criteria of



**Fig. 3.** Core data for drilling Sites U1589 (a-b) and U1592 (c-d). (a and c) Chronostratigraphic units, core lithologies and identified eruption units. Caldera-forming silicic eruptions are shown in red, and smaller eruptions in blue. The letters denote the newly recognized tephra A to L. (b and d) Age-depth curves based on biostratigraphic and tephrostratigraphic constraints. The red bars are previously known volcanic layers of known ages (Wulf et al., 2020; Druitt et al., 2024a). The biostratigraphic ages are taken from Druitt et al. (2024b); those <0.5 Ma are commonly older than tephrostratigraphic ages due to upward reworking associated with the very high sedimentation rates. Biostratigraphic ages >0.5 Ma are considered to be reliable as the associated sedimentation rates were lower. Magnetic reversal data are taken from Druitt et al. (2024b) and are in broad agreement with the biostratigraphic ages.

Rohling et al. (2015). Ages of previously unknown eruptions were estimated by extrapolating sedimentation rates between horizons of known age (Table S8). These age estimates are approximate due to the highly dynamic sedimentary environments of the rift axes, and they will be refined in a future publication integrating data from all the drill sites. The core chronostratigraphy was integrated with previously published seismic stratigraphy to create an integrated record of volcanism and tectonics (Fig. 2; Preine et al., 2025). See the Supplementary Text for full details of the methods used.

## 4. Results

### 4.1. Rift basin stratigraphy

Drilling penetrated the entire rift volcano-sedimentary fills and bottomed out in limestone/marble basement at both sites (593 mbsf [meters below sea floor] at U1589 and 510 mbsf at U1592; Figs. 2 and 3). The drilling recoveries above basement were 78 % (Hole A) and 92 % (Hole B) at Site U1589, and 71 % (Hole A) and 50 % (Hole B) at Site U1592 (Fig. S2). We divide the rift fills into five chronostratigraphic Units, CSK1A, CSK1B, CSK2, CSK3A and CSK3B from the top down, based on biostratigraphic, tephrostratigraphic, and seismic-stratigraphic data (Fig. 3; Tables 1 and S9). Unit CSK1A is dominated by lapilli ashes of silicic composition. Units CSK1B and CSK2 have upper parts dominated by oozes with many thin ashes of mafic to silicic composition and lower parts dominated by thick silicic lapilli ashes. Unit CSK3A consists of oozes and sands (dolomitic at Site U1592) with dispersed ash layers, and Unit CSK3B consists of siliciclastic sediments and bioclastic limestones. The basement at each site comprises limestones or marbles of Early Pleistocene to Cretaceous age.

The volcanic layers present in Units CSK1A to CSK3A range from beds <0.1 m thick to megabeds several tens (and in one case more than two hundred) of meters thick (Fig. 2; Figs. S3 to S7). By ‘megabed’ we mean an unusually thick and laterally extensive bed that differs in

composition from the host sediments (e.g., Sawyer et al., 2023; Druitt et al., 2024a). We attribute most of the volcanic layers to eruptions based on an abundance of glassy ash and lapilli, and to the chemical uniformity of vitric components (Figs. 4 and S8), while recognizing reworked volcanic layers by the presence of multiple vitric populations or abundant lithics.

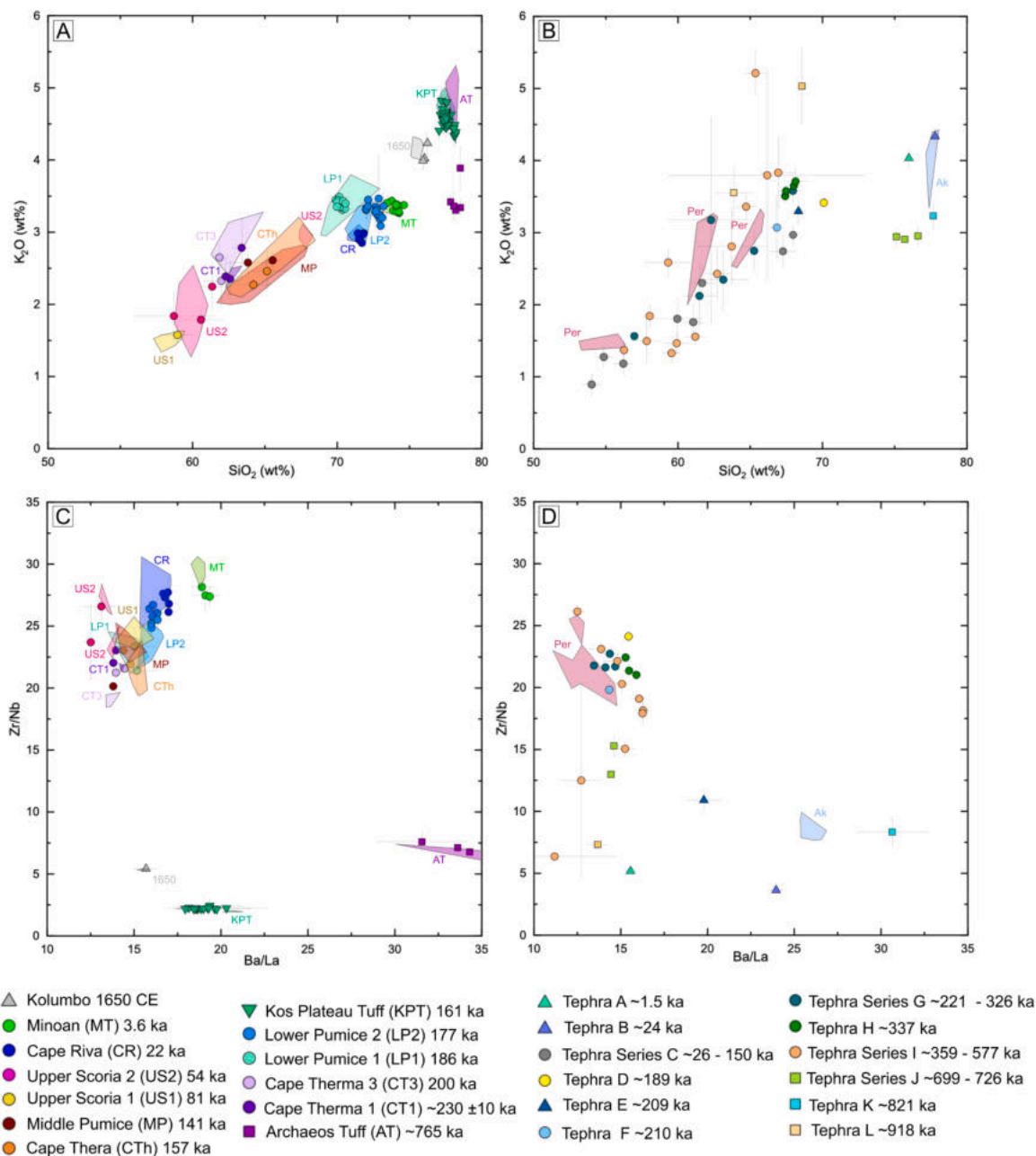
The products of most Santorini explosive eruptions known from onland studies (Fig. 1b) are recognized in the cores, based on chemical correlations (Tables 1 and S4; Figs. 2–4; Fig. S8). The Minoan (3.6 ka) and Cape Riva (22 ka) eruptions occur in chronostratigraphic Unit CSK1A, the Upper Scoria 2 (54 ka) to Lower Pumice 1 (186 ka) eruptions occur in Unit CSK1B, and the Cape Therma 1 (230 ± 10 ka) to Archaeos (~765 ka) eruptions occur in Unit CSK2. Tephra from three of the eruptions occur within sapropels (Upper Scoria 1 in Sapropel S3 at Site U1592; Lower Pumice 2 in Sapropel S6 and Cape Therma 3 in Sapropel S7 at Site U1589), as previously noted (Wulf et al. 2020). Biotite-bearing rhyolitic pumice from the 1650 CE eruption of Kolumbo Volcano occurs at the top of Unit CSK1A, and is the youngest volcanic layer in the cores. In both basins an extensive megabed of graded ash (up to 210 m at Site U1592) is present in Unit CSK1B between the Lower Pumice 2 (177 ka) and Cape Thera (157 ka) eruptions. This unit correlates chemically and mineralogically with the Kos Plateau Tuff (161 ka) from Kos caldera, 120 - 140 km to the east (Metcalfe et al., 2025). We show the four CFSEs of Modern Santorini, the Archaeos Tuff of Ancient Santorini, and the Kos Plateau Tuff in warm colors on Fig. 2e and f.

Forty previously undocumented tephra layers occur in the cores (Fig. 3; Tables 1 and S5), some of which we grouped into stratigraphically and compositionally related tephra series (Fig. 4). Tephra A is a biotite-bearing rhyolitic ash that has a trace element chemistry similar to that of the 1650 CE eruption of Kolumbo. Tephra B to H have the chemistries of magmas from Modern Santorini, with the exception of B and E, which derive from unknown sources, possibly Kolumbo. Tephra Series I is andesitic to dacitic; most of the ash layers compositionally overlap the fields for the onland Peristeria center (Modern Santorini),

**Table 1**  
Summary of rift volcanic stratigraphy.

Chronostratigraphic Unit	Depth U1589 mbsf	Depth U1592 mbsf	Composition	Age (ka)	Eruptions and ages*
CSK1A	0 – 53	0 – 15	Silicic lapilli ashes	0.4 – 26	Kolumbo, 0.4 ka Tephra A, 1.5 ka Minoan, 3.6 ka Cape Riva, 22 ka Tephra B, 24 ka
CSK1B	53 – 86	15 – 31	Oozes with multiple ashes	54 – 157	Upper Scoria 2, 54 ka Upper Scoria 1, 81 ka Middle Pumice, 141 ka Tephra series C, 26 - 150 ka Cape Thera, 157 ka Kos Plateau, 161 ka
	86 – 200	31 – 266	Silicic lapilli ashes	161 – 186	Lower Pumice 2, 177 ka Lower Pumice 1, 186 ka Tephra D, 189 ka
CSK2	200 – 302	266 – 348	Oozes with multiple ashes	200 – 725	Cape Therma 3, 200 ka Tephra E, 209 ka Tephra F, 210 ka Cape Therma 1, 230 ± 10 ka Tephra Series G, 221 - 326 ka Tephra H, 337 ka Tephra series I, 359 - 577 ka Tephra series J, 699 - 726 ka Archaeos, ~765 ka Tephra K, 821 ka Tephra L, 918 ka
CSK3A	302 – 313	348 – 402	Silicic lapilli ash	~765	
	313 – 375	402 – 490	Oozes with rare ashes (U1589)	~765 – 1600	
CSK3B	375 – 593	490 – 510	Siliciclasts (U1592) Siliciclastics (U1589)	1600 – 2000	
Basement	>593	>510	Bioclastic limestone (U1592) Limestone and marble	Eocene to Cretaceous (U1589) Early Pleistocene (U1592)	

\* Age estimates of previously unknown (lettered) tephra or tephra series are approximate



**Fig. 4.** Major and trace element discrimination diagrams for volcanic layers in the cores of Sites U1589 and U1592. (a and c) Glass compositions from the cores (data points) that correlate with previously known eruptions (colored fields showing the compositional ranges from previous studies, Kutterolf et al., 2021; Druitt et al., 2024a). Also shown are the known compositional ranges for Peristeria (Per) and Akrotiri (Ak) onland volcanics (new glass analyses in table S2). (b and d) Glass compositions from the cores for previously unknown eruptions or series of eruptions (Tephra layers or Tephra series A to L, with the data points for series enclosed by colored fields).

while some with higher K<sub>2</sub>O and lower Zr/Nb may derive from other, as yet undetermined, volcanic fields. Tephra Series J is rhyolitic with a chemical signature intermediate between Modern and Ancient Santorini, while Tephra K, also rhyolitic, chemically resembles the Archaeos Tuff. Tephra L is high in K<sub>2</sub>O, and may derive from a different volcanic field. With the exception of Tephra H (>4 m thick), all these layers are ≤42 cm thick, and are volumetrically much smaller than the TPF megabeds.

#### 4.2. Age-depth relationships

The age-depth curves obtained by integrating biostratigraphic ages with tephrostratigraphic ages for the previously dated volcanic units

reveal that the tephrostratigraphic ages are generally younger than the biostratigraphic ages in Units CSK1A, CSK1B and CSK2 (Fig. 3). We attribute this to upward reworking of microfossils in the highly dynamic rift environments. In what follows we use the tephrostratigraphic ages in Units CSK1A, CSK1B and CSK2 and the biostratigraphic ages in Unit CSK3A (no ages were obtained from Unit CSK3B). Estimated ages of Tephra A to L are given in Table 1. The base of Unit CSK3A has an age of ~1.6 Ma; by extrapolation, we estimate that the base of Unit CSK3B at Site U1592 would be ~2 Ma and at Site U1589 (where CSK3B is thicker) would be >2 Ma (Fig. 3). Biostratigraphic dating of cores from Site U1599 (Fig. 1a) provides an even earlier constraint on the onset of rifting of the Anafi Basin (>2.7 Ma; Druitt et al. 2024b).

### 4.3. Core-seismic correlation

Correlation of the core chronostratigraphy with the published seismic stratigraphy reveals a consistent seismo-stratigraphy of the Anhydros and Anafi Basins (Fig. 2). Preine et al. (2022a,b) recognized six seismic stratigraphic packages in each of the two basins. They also recognized two major seismic reflection onlap surfaces: one (their onlap surface h4) that is particularly prominent in the Anhydros Basin and the other (their onlap surface h6) that is prominent in the Anafi Basin. They interpreted these onlaps as recording phases of rapid subsidence with relatively little sedimentation ('rift pulses'). The IODP Exp. 398 cores now make it possible to redefine the previously only visual correlation between the two basins (Table S9), such that the two major onlap surfaces previously identified are in fact the same (shown as a solid red line on Fig. 2a-b). This single major onlap surface directly underlies Unit CSK1B; it is formed of inclined packages of marine sediments and minor tephra layers of Unit CSK2, and is directly overlain by thick packages of CFSE megabeds of the TPF and Kos Plateau Tuff eruptions.

### 4.4. Sedimentation rates

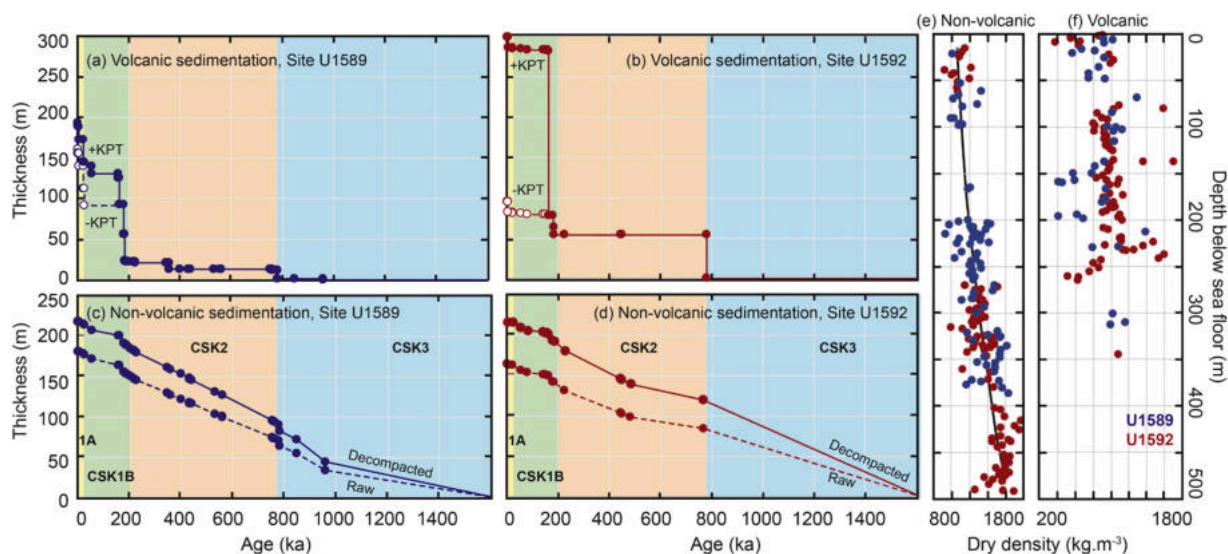
The age-depth curves show upward-increasing sedimentation rates from Unit CSK3A to Unit CSK1A, with prominent inflections at the bases of Units CSK2 and CSK1B (Fig. 3). Taking real (variably compacted) thicknesses, 380 m of sediment have accumulated in the Anhydros Basin, and 480 m in the Anafi Basin, over the last 1.8 Ma. Separating the volcanic and non-volcanic inputs (neglecting any dispersed volcanic contribution in the background intervals), and decompacting the non-volcanic sediment using shipboard bulk density data (Fig. 5e-f), the sedimentation rates for non-volcanic sediment are 9–14 cm kyr<sup>-1</sup> in Unit CSK3A and 12–18 cm kyr<sup>-1</sup> in Units CSK2 to CSK1A, reflecting the similar sedimentation histories of the two basins (Fig. 5c-d). These rates of background marine sedimentation resemble those measured in shallow gravity cores in the Aegean Sea surrounding, and east of, Santorini (averaging about 3–10 cm kyr<sup>-1</sup>; Kutterolf et al., 2021). The time-averaged volcanic sedimentation rates in the two basins, essentially zero until 765 ka, are 25–40 cm kyr<sup>-1</sup> from 765 to 0 ka and 90–140 cm kyr<sup>-1</sup> over the last 186 kyr (Fig. 5a-b).

## 5. Discussion

### 5.1. Eruption-fed megabeds in a volcanic rift system

While the rift basins contain numerous cm-thin tephra layers of probable fallout origin, several of the thickest volcanic layers (Figs. 2 e-f, S7) can be classified as volcanoclastic megabeds (Druitt et al., 2024b; Metcalfe et al., 2025). Based on chemical homogeneity, paucity of strong internal reflections, and lack of internal bioturbation, we interpret each megabed as having been emplaced rapidly by an eruption or shortly afterwards by remobilization of a primary deposit. This rapid deposition could have occurred by submarine gravity flows, fallout through the water column, and sinking of pumice rafts (Freundt et al., 2023). Megabeds thicker than a few meters are interpreted as the deposits from pyroclastic currents (from either subaerial or submarine vents) that entered the water column and, by entraining sea water, transformed into water-supported gravity flows (Clare et al., 2023; Cas et al., 2024). This interpretation is based on the fact that many of the Santorini eruptions that generated submarine megabeds in our cores are known from onland studies to have discharged voluminous pyroclastic currents into the sea (Druitt et al. 1999), as did the Kos Plateau Tuff eruption (Allen, 2001). The thickest megabeds are the Kos Plateau Tuff (210 m) and the Archaeos Tuff (~50 m) at Site U1592, and Lower Pumice 2 (38 m) at Site U1589 (Fig. 2 e-f). The Archaeos and Kos Plateau megabeds have been described in detail by Druitt et al. (2024a) and Metcalfe et al. (2025), respectively. Since the climactic phases of CFSEs typically last a few days or less, perhaps followed by several years to decades of remobilization, emplacement of each megabed was instantaneous on a geological timescale.

The prominence of megabeds in the SATZ basins is in marked contrast with sedimentation in most non-volcanic rifts. Although major re-sedimentation events due to flood input, slope instability or tsunami passage can produce megabeds, most sediment accumulation in non-volcanic rifts occurs incrementally (e.g., Gawthorpe et al., 2022). Many volcanic rifts are characterized by lava effusion and high supply rates of volcanoclastic material (Hutchison et al., 2016; Chen et al., 2022). The SATZ is an example of a rift system hosting CFSE volcanoes that generate huge, eruption-fed submarine deposits (Cas et al., 2024).



**Fig. 5.** Sedimentation data from rift Sites U1589 and U1592 from 1.6 Ma to the present day. (a-b) Cumulative thicknesses of volcanic layers versus age. The data are shown including and excluding the Kos Plateau Tuff (KPT), which does not derive from Santorini-Kolumbo. (c-d) Cumulative thicknesses of non-volcanic sediments versus age, both raw and decompacted. (e-f) Sediment dry density data from shipboard moisture and density measurement database (Druitt et al., 2024b). Non-volcanic sediments compact with burial, whereas volcanic products do not. The black line on (e) is a best fit polynomial to the non-volcanic sediment data (both sites combined) for dry bulk density ( $\text{kg m}^{-3}$ ),  $0.0015D^2 + 0.3965D + 1059.3$ , where  $D$  is depth in meters, that was used to decompact the non-volcanic sediment thicknesses in (c-d).

Such gravity flows can travel over a hundred km from source and re-mould the seafloor landscape by erosion and by infilling of bathymetric troughs on an arc-wide scale, as shown by the Kos Plateau Tuff ash megabed in the Anafi Basin (Metcalfe et al., 2025).

5.2. Rift subsidence history

The new data allow us to revisit the published subsidence models for the Anhydros and Anafi basins based on seismic stratigraphy (Preine et al., 2022a). The sedimentation rates and stratigraphies of the two rift basins show that they have had similar histories of development since the onset of subsidence >2 Ma (Figs. 2 and 3). Using the thicknesses of sediment in each basin, the time-averaged subsidence rates since 1.8 Ma in the Anhydros and Anafi Basins have been 21 cm kyr<sup>-1</sup> and 27 cm kyr<sup>-1</sup>, respectively (the higher rate of the Anafi Basin possibly being due to its greater thickness of volcanic megabeds (Fig. S9), lower elevation, and larger size; Nomikou et al., 2018). However, the subsidence may have been non-uniform, with a phase of accelerated rifting recorded by the seismic reflection onlap surface (Fig. 3a-b). Preine et al. (2022a) proposed that this onlap corresponds to a phase of rapid extension during which sedimentation failed to keep up with subsidence. The IODP data support this interpretation by showing that following the extension

event the rift basins were deep enough to accommodate a thick sequence of rapidly emplaced CFSE megabeds. This is particularly the case in the Anafi Basin where 235 m of ash was emplaced over 25 kyr (from 186 to 161 ka) and 210 m of ash from the Kos Plateau Tuff eruption was deposited in a single, geologically instantaneous event (Metcalfe et al., 2025). The Anafi basin must therefore have been >200 m deep, and the Anhydros Basin >100 m deep, to accommodate these megabeds. We therefore infer that during extension the basins became starved of sediment and subsequently acted as depocenters for eruption-fed gravity flows. These conclusions remain valid even when we take into account glacio-eustatic sea level changes, with minima of -80 to -120 m over the last 500 kyr (Grant et al., 2014).

We summarize the development of the Anhydros and Anafi Basins in three stages (Fig. 6): **Stage 1** (*Sedimentation ≈ subsidence*): following the onset of rifting at >2 Ma, the sedimentation of Units CSK3 and CSK2 probably approximately kept up with subsidence, although transient fluctuations were likely. **Stage 2** (*Sedimentation < subsidence*): the rifting rate accelerated, and the basins deepened because sedimentation of Unit CSK2 could no longer keep up with subsidence. **Stage 3** (*Sedimentation > subsidence*): the deep basins formed during Stage 2 acted as depocenters into which pyroclastic currents of the Santorini TPF and Kos CFSEs poured, forming thick volcanic megabeds (Units CSK1B and

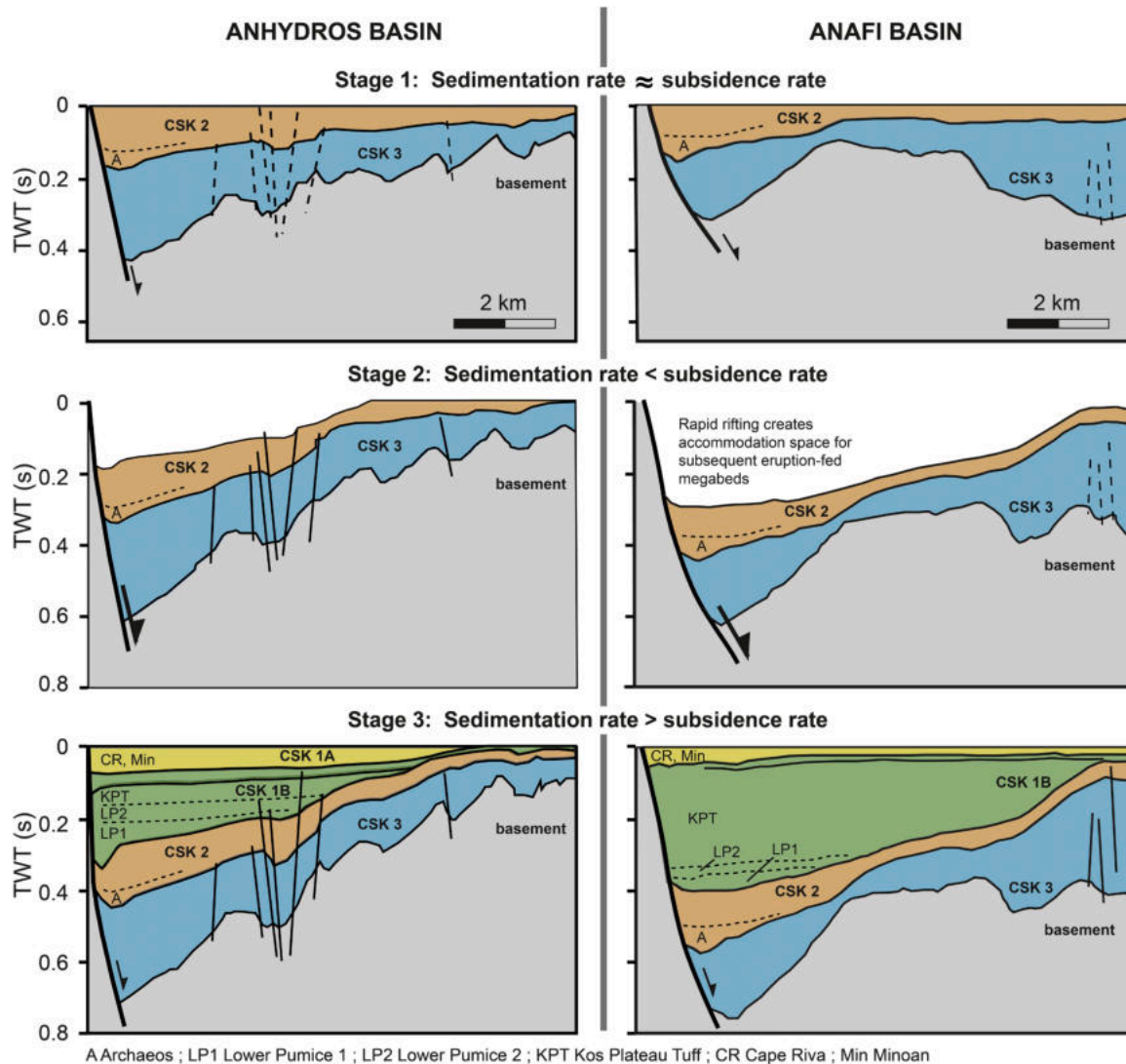


Fig. 6. Three stages in the development of the Anhydros and Anafi Basins described in the text, reconstructed from seismic sections and the drill core analysis. The reconstructions are based on successively flattening the top of each unit, as in Preine et al. (2022a). The colors of the chronostratigraphic units are the same as in Figs. 2c-d and 3a-c.

CSK1A). The seismic reflection onlap surface is interpreted as the combined product of (1) the existence of deep, rapidly subsiding basins, and (2) a change in sedimentation style from normal marine sediments (and thin tephra layers) to large, bathymetry-filling gravity flows. Subsidence rate during Stage 3 was insufficient to strongly tilt the CFSE megabeds of the Anhydros Basin, the top surfaces of which remain sub-horizontal (Fig. 2). However, updrag of Unit CSK1B against the master fault of the Anafi half-graben (Fig. 2d) shows that significant subsidence continued in that basin during Stage 3 (Preine et al. 2022a).

Since sedimentation rate in the rifts increased with time during emplacement of Units CSK3A and CSK2 (Figs. 3 and 5), basin deepening during Stage 2 must have been due to increased subsidence rate, not decreased sedimentation rate. The duration of Stage 2 is not well constrained. Preine et al. (2022a) presented arguments based on the nature of internal seismic reflections that the transition from Stage 1 to Stage 2 was rapid, and that Stage 2 involved a well-defined, transient increase in extension rate or 'rift pulse' that possibly lasted several tens of thousands of years prior to the onset of caldera-forming eruptions of the TPF at ~250 ka. However, a more gradual acceleration of extension rate from Stage 1 to Stage 2 cannot be ruled out. In what follows we refer to Stage 2 as a phase of rapid rifting, 'rapid' meaning relative to local sedimentation rate and not necessarily relative to rifts globally. Precise subsidence-time curves for the Anhydros and Anafi Basins await full structural re-evaluation of the rift basins in the light of IODP core-seismic integration.

### 5.3. Volcanic history of Santorini recorded by the rift successions

The results confirm the explosive volcanic record of Santorini known from onland studies and provide new insights.

First, the earliest tephra layer with Ancient Santorini chemistry at our rift sites is dated at 821 ka (Tephra K), older than the earliest volcanics dated onland (~650 ka; Druitt et al., 1999). This confirms published assertions that the onset of SATZ rifting (>2 Ma) preceded volcanism at Santorini and is consistent with the presence of an ancient 'proto-Anhydros' Basin buried beneath the volcano (Heath et al., 2019; Preine et al., 2022a). However, the presence of an even earlier tephra layer (Tephra L; 918 ka) with a high-K signature, as well as a 2.4-m-thick pumice layer ~1.6 Myr old recovered at IODP Expedition 398 Site U1600 (Fig. 1a; Druitt et al., 2024b), shows that one or more other volcanic centres already existed in the area before the eruption of Tephra K. It is also possible that Christiana Volcano, which was active until ~1.6 Ma (Preine et al., 2022a), partly predated opening of the SATZ.

Second, the earliest products of Modern Santorini (Tephra Series I) date from about 577 ka, broadly consistent with onland dating (530 ka; Druitt et al., 1999). A package of rhyolitic tephra layers with trace element signatures intermediate between those of Ancient and Modern Santorini (Tephra Series J) is dated 699 – 726 ka.

Third, the magma output rate and frequency of highly explosive eruptions during the TPF period exceeded those of the first ~300 kyr years of Modern Santorini's history, as reflected in the great thickness of basin-hosted megabeds from the TPF eruptions (Fig. 5a-b). Between 30 and 50 % of the basin fills accumulated in just one tenth of the lifetimes of the rifts.

### 5.4. Rift modulation of caldera volcanism?

We now focus on the possible relationships of lithospheric rifting to the transition to caldera-forming silicic eruption (CFSE) activity at Modern Santorini ~250 kyr ago. By ground-truthing the seismic stratigraphy of Preine et al. (2022a), and integrating volcanic, sedimentary and tectonic histories (Fig. 7), we have shown the existence of a single major seismic reflection onlap surface and have confirmed that the majority of the TPF eruptions (and the Kos Plateau Tuff) lie directly above that onlap. We have also shown that these deposits (Units CSK1A

and CSK1B) consist largely of stacked gravity-flow-emplaced megabeds, representing a marked change in sedimentation mechanism from the nonvolcanic sediments and lesser tephra layers (Unit CSK2) beneath the onlap surface. This opens up two possible interpretations of the relationship between rifting and CFSE volcanism.

**Model 1.** In this interpretation the transition of Modern Santorini to CFSE activity was driven purely by internal magmatic processes and was unrelated to tectonics. Subsidence of Stage 2 progressively deepened the rift basins until they formed bathymetric troughs. Then, when Santorini became highly explosive, the basins simply acted as passive receptacles for the pyroclastic gravity flows, whose thick deposits onlapped the basin margins in a manner analogous to a terrestrial pyroclastic flow filling a valley (e.g., Cas et al. 2024).

**Model 2.** In this interpretation the transition to CFSE activity was the result of prolonged extension and a phase of rapid subsidence and basin deepening (Stage 2; Preine et al. 2022a), with a direct relationship between rifting and explosive volcanism. The question arises in Model 2 as to whether rifting caused the CFSE volcanism or whether an increased crustal magmatic activity prior to the onset of the TPF triggered the phase of rapid rifting by thermally weakening the lithosphere, lubricating faults, and accommodating strain (Brune et al., 2023). Since Plio-Quaternary rifting in the central and southeastern Aegean is regionally driven (Jolivet et al., 2013; Sternai et al., 2014), we favour that in Model 2 the elevated levels of crustal magmatism that drove the TPF were the result of rifting, not vice versa.

While we cannot definitively exclude Model 1, the location of Santorini in the western Anhydros Basin, the evidence for rapid rift extension, and the well documented influence of tectonics on volcanism in the area (Kokkalis and Aydin, 2012; Hooff et al., 2017; Heath et al., 2019, 2021; Preine et al. 2022a; Drymoni et al., 2022) lend support to Model 2. We therefore explore three mechanisms by which rifting may have modulated the transition of Modern Santorini to CFSE activity (Fig. 8).

**Change in mantle melt extraction.** Rapid rifting prior to ~250 ka may have increased the flux of mantle-derived melts, fluids and heat to the transcrustal magmatic system. Two observations support a change in mantle supply to Modern Santorini at this time. First, the increased eruption rate of the TPF compared to the preceding period (Fig. 1b) implies an increased flux of basalt into the crust. Increased mantle flux was necessary to provide the extra mass and heat to the transcrustal magmatic system to generate the >120 km<sup>3</sup> of chemically evolved melts erupted as the TPF, an output possibly several times higher than that during the preceding Peristeria period (Figs. 1b and 5 a-b). Second, magmas of the TPF are poorer in incompatible elements such as K, Rb, Th and Zr than those of the preceding Peristeria Volcano, suggesting a change in the nature of the incoming parental melts (Fig. S10; Huijsmans et al., 1988; Druitt et al., 1999). One explanation for increased mantle supply might have been an increase in the degree of mantle melting due to lithospheric thinning. However, slowly extending rifts like the SATZ (<few mm yr<sup>-1</sup>; Feuillet, 2013) do not generate large amounts of decompression melting in thermomechanical models at the mantle potential temperatures beneath Santorini (1350 – 1440°C; Flaherty et al., 2022) (see figure 9 in Fioraso et al., 2024). Several kilometers of localized, extension-driven mantle upwelling beneath the volcano would therefore have been necessary to generate even a few % of additional melting (Sternai, 2020). More likely, transient changes in asthenospheric flow rates and patterns accompanying rifting increased the rate of melt extraction from more refractory, higher-temperature regions of the mantle wedge (e.g., Leeman 2020), supplementing the input from more enriched, fluid-fluxed regions and raising the total supply of basalt (e.g., Hutchison et al., 2016).

**Increased melt percolation through the transcrustal magmatic system.** Rapid rifting may also have caused an increase in crustal shearing, rock permeability, shear-induced melt segregation, and gravitational instabilities in magmatic mushes of the transcrustal magmatic system (Kohlstedt and Holtzman, 2009; Sparks et al., 2019; Weinberg et al., 2021). Given favorable stress conditions, these effects may have

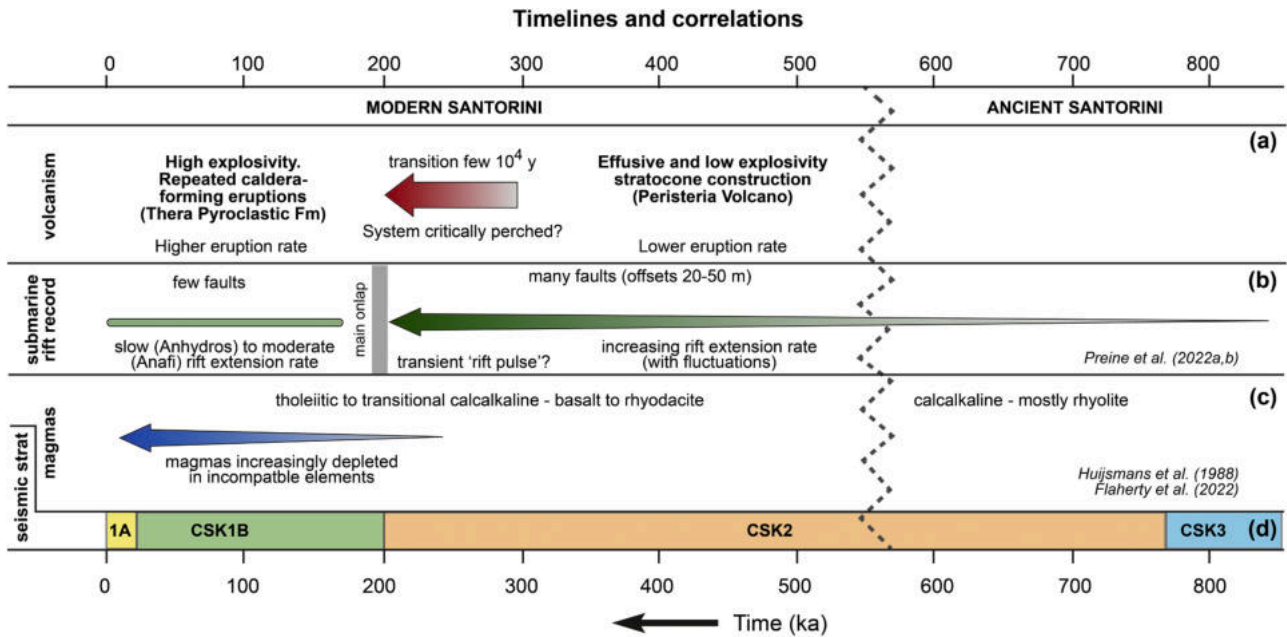


Fig. 7. Schematic summary of major volcanic and tectonic events. (a) Volcanic activity. (b) Rifting record from core-seismic integration. (c) Geochemical features of Santorini magmas from previous studies. (d) Chronostratigraphic stratigraphy in the submarine cores. Source references are given on the figure.

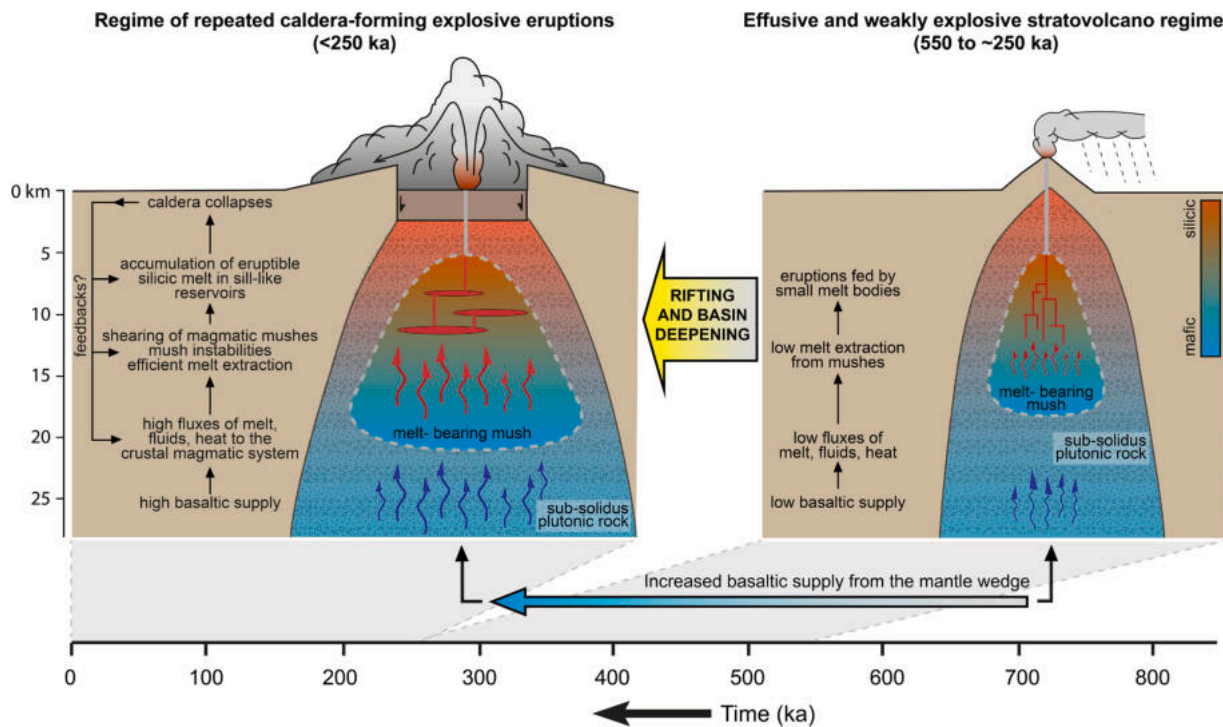


Fig. 8. Schematic (Model 2) for the modulation of volcanic activity at Modern Santorini by the phase of accelerated rifting that took place between about 300 and 250 ka. Subsidius plutonic rock of the transcrustal magmatic system is shown patterned in pale colours and melt storage regions are shown in stronger colours, with blue for mafic compositions and red for silicic compositions. The rifting is interpreted as having increased the fluxes of melts, fluids and heat from the mantle and through the transcrustal magmatic system, driving a transition from predominantly effusive and weakly explosive andesitic behavior to strongly explosive activity with repeated caldera-forming silicic eruptions. The blue arrows depict ascent of parental basalt, and the red arrows depict movement of silicic melts through the transcrustal crystal mushes.

increased the upward flow of silicic/intermediate melts and fluids through the transcrustal magma reservoir of Santorini, promoting formation of the large upper crustal lenses of gas-charged eruptible melt necessary to feed CPSEs (Flaherty et al., 2018).

**Changes in magma storage architecture.** A third possibility is that

prolonged rifting and basin deepening caused a change in the magma reservoir architecture of the transcrustal magmatic system. Stresses induced by deepening volcanic rifts influence magma storage architecture by changing the balance between regional and unloading stress components (Ferrante et al., 2024). As a rift deepens, the minimum

principal stress can rotate towards the vertical, promoting development of complexes of stacked sills in the middle to lower crust. Top-down crustal stress changes due to progressive deepening of the western Anhydros Basin may therefore have favored a transition at Santorini from dyke-dominated magma ascent to the development of stacked sills that could then have trapped the incoming basalt and led to thermal runaway. Repeated transient unloading due to caldera collapses could then have maintained a sill-dominated architecture through feedback effects, while also maintaining high melt fluxes through the magmatic system, possibly explaining the continuation of CFSE-dominated activity into the Holocene (Corbi et al. 2015).

In summary, Model 2 envisages that a rifting-induced increase in flow of melts, fluids and heat into, and through, the transcrustal magmatic system, coupled with changes in magma reservoir architecture, may have amplified the internal magmatic processes at Santorini and driven it ~250 kyr ago from an andesitic stratovolcano into a regime of repeated CFSE activity (Fig. 8). The effects of tectonic forcing would have been facilitated if, following over 300 kyr of activity of Modern Santorini from 550 to 250 ka (plus the additional ~300 kyr of Ancient Santorini), the transcrustal magmatic system was already hot and primed for thermomechanical runaway. If the phase of rapid rifting (Stage 2 of Fig. 6) lasted at least several tens of thousands of years, then this time was available for the transcrustal magmatic system to react to it before the first CFSE of Modern Santorini at 186 ka (Figs. 1b and 7).

Distinguishing definitively between Models 1 and 2 will require detailed re-evaluation of the basin evolution in the light of the new core-seismic integration. However, a correlation of eruptive events across the South Aegean Volcanic Arc provides additional weight in favour of Model 2, as now described.

### 5.5. A flare-up of caldera-forming activity on the South Aegean Volcanic Arc

The western sector of the volcanic arc is characterized by small, commonly monogenetic centers, whereas the central (Christiania–Santorini–Kolumbo) and eastern (Kos–Nisyros–Yali) sectors are both characterized by composite volcanoes with calderas (Vougioukalakis et al., 2019). This suggests that higher mantle supply rates, and/or crustal stress conditions, are more favourable to the formation of calderas in the central and eastern sectors than further west. Interestingly, the two caldera systems – Santorini and Kos – lie in mature rift systems, reinforcing the relationship between rifting and explosive activity. The rift system hosting the Kos caldera is oriented WSW–ENE and, like the SATZ, is of late Pliocene to Pleistocene age (Papanikolaou et al., 2018).

It is notable that the Kos Plateau Tuff eruption, the largest eruption of the Kos volcanic centre since the ~500 ka Kefalos eruptions (Bachmann et al., 2019), occurred at 161 ka just after the Lower Pumice pair of CFSEs (186 and 177 ka) at Santorini. Moreover, the Kos caldera has remained in an explosive state to the present day, with at least two >10 km<sup>3</sup> explosive eruptions from Nisyros and Yali Volcanoes on the rim of the Kos caldera (Kutterolf et al., 2021). This raises the possibility of broad synchronizations between the rift-hosted Santorini and Kos magmatic systems, and a 0.2 Ma flare-up of caldera-forming volcanism in the central and eastern sectors of the volcanic arc. The connection is likely to have arisen from the same plate-wide stresses driving rifting at both systems (Jolivet et al., 2013), and from the Kos–Nisyros–Yali magmatic system having interacted with its rift environment in a similar manner to Modern Santorini, as inherent in Model 2 above.

An ignimbrite flare-up is a period of abnormally high explosive activity along a segment of volcanic arc or continental rift, such as the Central Volcanic Zone of the Andes (de Silva et al., 2015), the Taupō Volcanic Zone of New Zealand (Gravely et al., 2016), or the East African Rift (Hutchison et al., 2016). Flare-ups occur on a range of timescales, from 10<sup>4</sup> to ≥10<sup>7</sup> yrs, and can explosively discharge up to 10<sup>3</sup>–10<sup>4</sup> km<sup>3</sup> of silicic magma (Gravely et al., 2016). The CFSE flare-up at Santorini and Kos–Nisyros–Yali volcanic fields is small relative to other examples

on a global scale in terms of timescales (a few 10<sup>4</sup> yr) and total (Santorini TPF plus Kos) volumes of erupted magma (~200 km<sup>3</sup>). However, it is similar in spatial, temporal and volume scales to that which took place in the rift valley of Ethiopia in the Middle Pleistocene (Hutchison et al., 2016). Our integrated core-seismic dataset suggests that the flare-up at Santorini and Kos may have been triggered by a phase of accelerated, regional rifting, and that the magmatic systems took at least several tens of thousands of years to react to the changes in lithospheric stresses and mantle supply. These findings show that the well-known multiscale nature of volcanic arc tempos (de Silva et al., 2015), and the effects of tectonic forcing, may extend down to the small temporal (10<sup>4</sup>–10<sup>5</sup> yr) and volume (10<sup>2</sup> km<sup>3</sup>) scales typical of individual caldera cycles at arc volcanoes (Bouvet de Maisonneuve et al., 2021).

## 6. Conclusions

1. The close association of caldera volcanism with lithospheric rifting in the central and eastern sectors of the South Aegean Volcanic Arc makes the region a natural laboratory for investigating volcano-tectonic couplings at calderas.
2. IODP Expedition 398 deep-drilled the volcano-sedimentary fills of the rift basins that lie northeast of, and underneath, Santorini (Santorini–Amorgos Tectonic Zone), ground-truthing existing seismic stratigraphy and generating a unique, high-resolution, core-seismic integration record of volcanic activity and lithospheric extension on a volcanic arc. The results reveal a possible example of rift modulation of an arc magmatic system on the 10<sup>4</sup>–10<sup>5</sup> yr timescales typical of caldera cycles.
3. Regionally-driven rift extension began >2 Ma, then accelerated into a rapid phase during which the marine basins deepened because subsidence outpaced sedimentation. These bathymetric troughs then subsequently acted as depocenters for thick sequences of lapilli-ash megabeds fed by caldera-forming eruptions of Santorini (Thera Pyroclastic Formation) and the Kos Plateau Tuff eruption of Kos Volcano.
4. The phase of accelerated rifting may have pushed Modern Santorini ~250 ka from a regime of effusive and weakly explosive behavior typical of andesitic stratovolcanoes into its present-day highly explosive, caldera-forming state. The rifting may have amplified the normal internal processes and feedbacks of magma ascent, storage and eruption by (i) increasing the flow of magmatic melts, fluids and heat through the transcrustal magmatic system, (ii) modifying the architecture of magma storage, and (iii) driving the system into thermomechanical runaway. Repeated transient unloading of the volcanic plumbing system by caldera collapses may have helped sustain the highly explosive activity into the Holocene.
5. Broadly simultaneous transitions of the rift-hosted Santorini and Kos volcanic fields into caldera-forming activity about 0.2 million years ago suggest that the eruptive states of the two magmatic systems are linked by plate-scale lithospheric stresses.

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## CRedit authorship contribution statement

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Writing – review & editing, Data curation. **Steffen Kutterolf**: Writing – review & editing, Data curation, Conceptualization. **Jonas Preine**: Writing – review & editing, Data curation, Conceptualization. **Sarah Beethe**: Writing – review & editing, Data curation. **Axel Schmitt**: Writing – review & editing, Data curation. **Christian Hübscher**: Writing – review & editing, Data curation, Conceptualization. **Paraskevi Nomikou**: Writing – review & editing, Data curation, Conceptualization. **Thomas A. Ronge**: Writing – review & editing, Project administration. **Carole Berthod**: Data curation. **Hehe Chen**: Data curation. **Shun Chiyonobu**: Data curation. **Acacia Clark**: Data curation. **Susan DeBari**: Data curation. **Ralf Gertisser**: Writing – review & editing, Data curation. **Raymond Johnston**: Data curation. **Olga Koukousioura**: Data curation. **Michael Manga**: Writing – review & editing, Data curation. **Molly McCanta**: Data curation. **Iona McIntosh**: Data curation. **Ally Peccia**: Data curation. **Masako Tominaga**: Data curation. **Yuzuru Yamamoto**: Data curation. **Adam Woodhouse**: Data curation. **Alexis Bernard**: Data curation. **Tatiana Fernandez Perez**: Data curation. **Christopher K. Jones**: Data curation. **Kumar Batuk Joshi**: Data curation. **Günther Kletetschka**: Writing – review & editing, Data curation. **Antony Morris**: Data curation. **Paraskevi Polymenakou**: Data curation. **Xiaohui Li**: Data curation. **Dimitrios Papanikolaou**: Funding acquisition. **David Pyle**: Writing – review & editing. **Pietro Sternai**: Writing – review & editing.

## Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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## Supplementary materials

Supplementary material associated with this article can be found, in the online version, at [doi:10.1016/j.epsl.2025.119633](https://doi.org/10.1016/j.epsl.2025.119633).

## Data availability

High-resolution seismic profiles from cruise POS538 can be accessed from Pangea at <https://doi.org/10.1594/PANGAEA.956579>. A selection of vintage seismic profiles can be found in the marine geoscience data system (MGDS) at <https://doi.org/10.26022/IEDA/327525> and <https://doi.org/10.26022/IEDA/331028>.

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